## The value of Doppler LiDAR systems to monitor turbulence, dust and ash intensity in order to enhance aviation safety in Iceland



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## Summary

In this study, we use a ground-based LiDAR (Fig. 1) system stationed near Reykjavik Domestic Airport, Iceland (Fig. 2), to evaluate turbulence intensity, dust and ash concentrations. For this purpose, we retrieved radial wind velocity (Fig. 3) to compute eddy dissipation (EDR) and backscatter signals for aerosol mass concentrations. The method was used to monitor and characterize a storm event in fall 2016 (Fig. 7). The preliminary results (Fig. 8) have been compared with the method using the vertical scan (O'Connor et al., 2010), and it reveals that the LiDAR observations can detect and quantify atmospheric turbulence with high spatial and temporal resolution.

## LiDAR Specifications

The Leosphere WINDCUBE 200s LiDAR

Wavelength: 1.54 µm Maximum range: 12 km Azimuth angle: 0 ~ 360°

Elevation angle: -10 ~ 190°



Fig 1. mobile LiDAR at IMO, Reykjavik



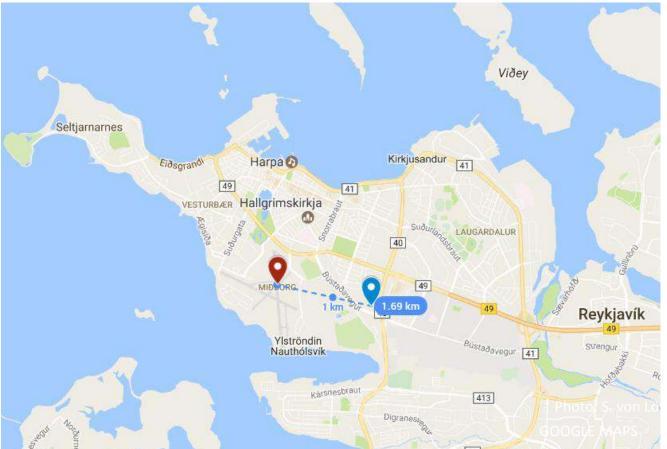


Fig 2. Location of Reykjavik (Left), the location of IMO (right, blue pin) and Reykjavik Domestic Airport (right, red pin).

## LiDAR set up

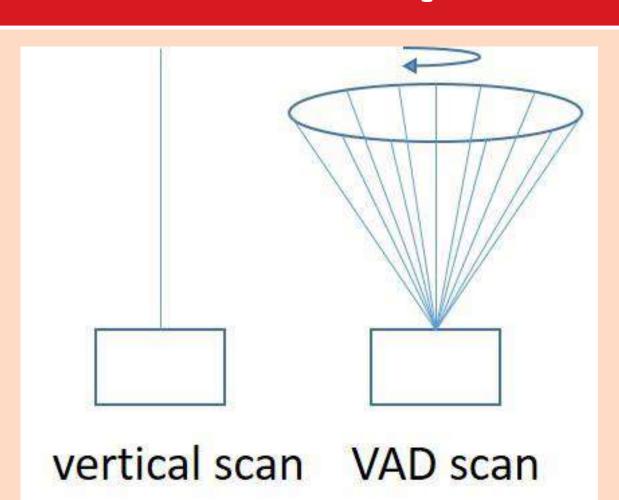


Fig 3. Two scan modes of LiDAR

For routine monitoring of turbulences the LiDAR system was programmed to do the following scans (Fig. 3):

- Continuous vertical scans
- 2 VAD scans four times an hour at 75° and 15° elevation angle, respectively
- Periodic technical scans during night

VAD Scan gets more information about horizontal wind component but requires homogeneous atmosphere.

## Kolmogorov Theory

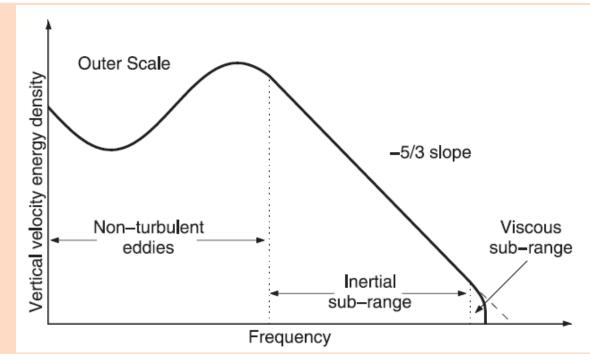
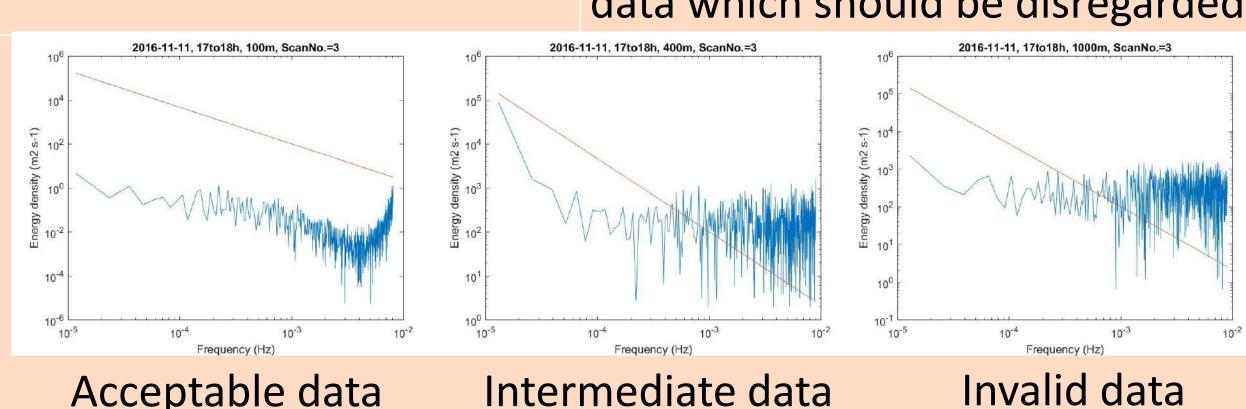


Fig 4. Schematic of vertical velocity energy density spectra vs frequency conforming to Kolmogorov's hypothesis (O'Connor et al., 2010)

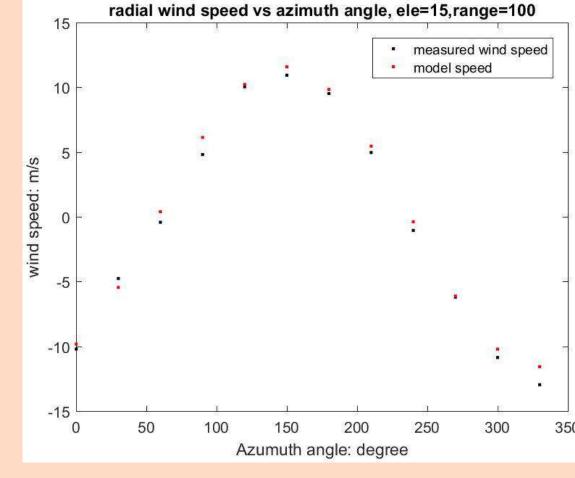
The Kolmogorov Theory (see the algorithm section below) can be applied when the atmosphere is isotropic, i.e. the energy spectrum agrees with -5/3 slope (Fig. 4). Fig. 5 shows some examples. There are acceptable results where the theory can be applied directly and invalid data which should be disregarded.



Acceptable data Fig 5. Some examples of energy spectrum

Invalid data

## Algorithm to compute wind speed



Measured and modelled wind velocity, the difference is fluctuation

EDR can be computed with the following algorithm:

- $D_{\nu} = C_{\nu} \epsilon^{2/3} s^{2/3}$
- Kolmogorov constant  $C_v \approx 2$
- $D_{v}(s) =$  $\langle [v'(r,\varphi,\theta)-v'(r,\varphi+\Delta\varphi,\theta)]^2 \rangle$
- $v'(r, \varphi, \theta)$  are the fluctuations Fluctuations are the difference between measured and modelled wind speed (Fig. 6).

# Application to real weather condition

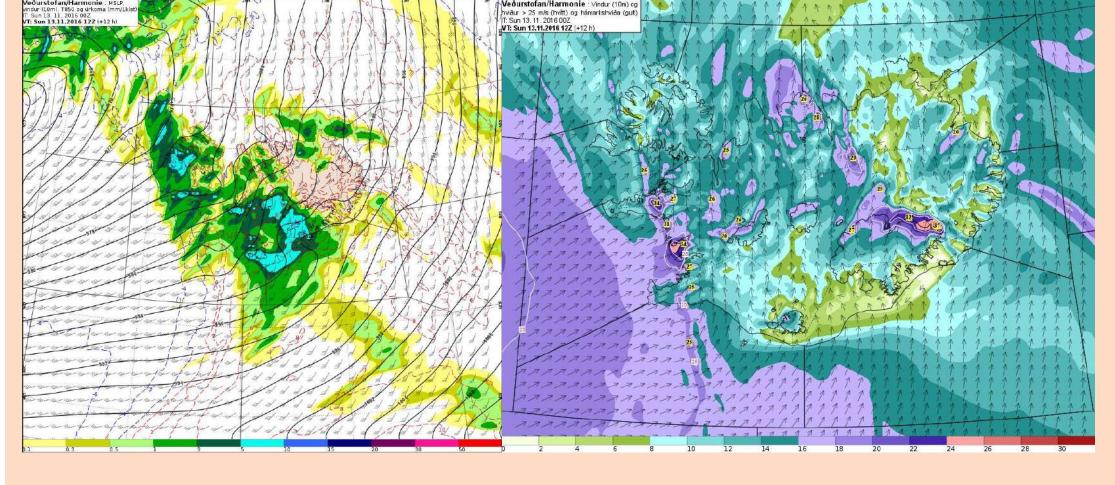
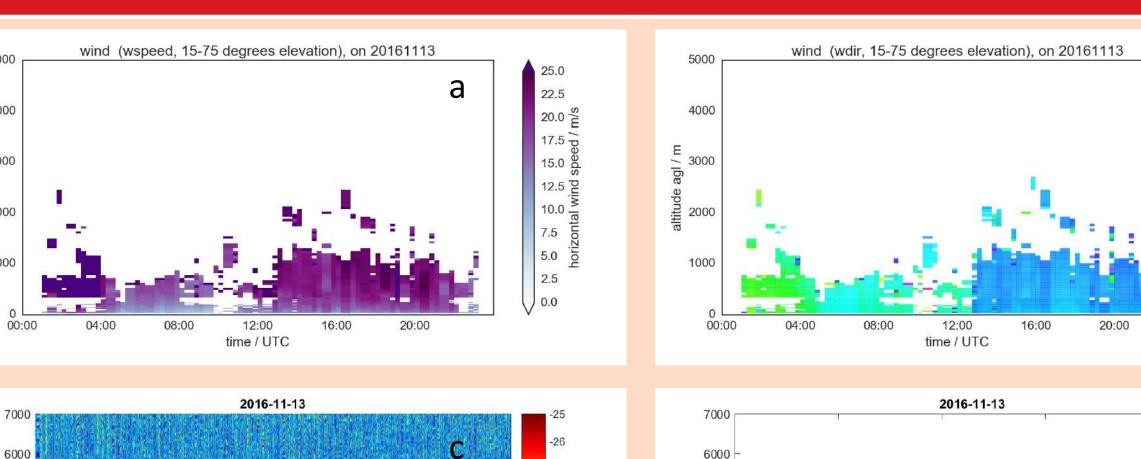
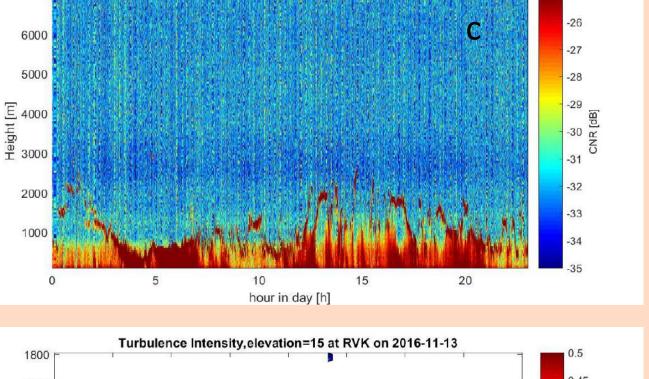


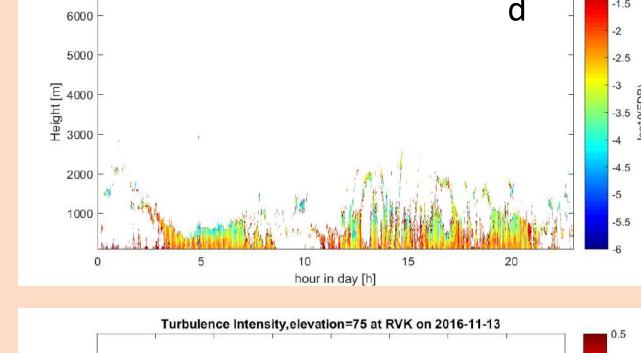
Fig 7. Weather forecast on 13 November 2016

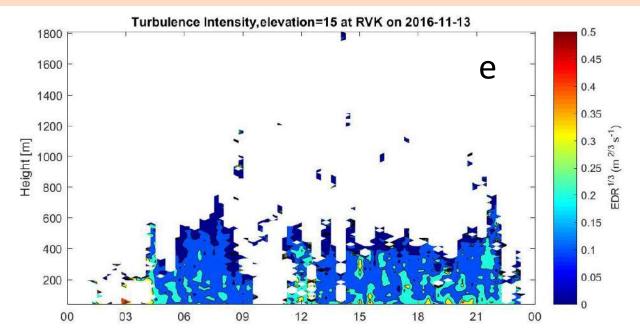
The algorithm was applied to a selected storm event, 13 November 2016, shown in Fig. 7. Resulting wind speeds and EDR is shown in Fig. 8.

### Results









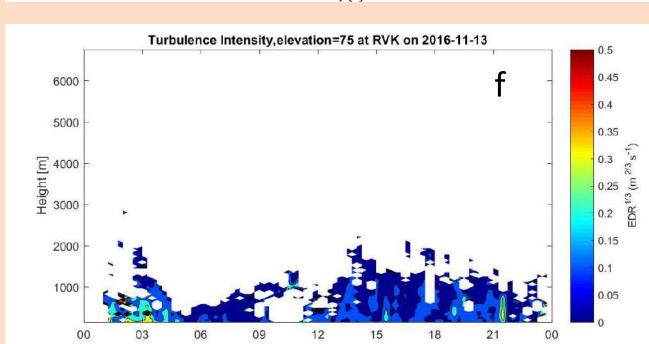


Fig 8. (a) Horizontal wind speed, (b) wind direction, (c) carrier-noise ratio, (d) eddy dissipation rate retrieved by vertical velocity (O'Connor et al., 2010), (e) eddy dissipation rate retrieved by 15° and (f) 75° degree VAD scan.

### Conclusion

Fig. 8 shows that wind information can be obtained by LiDAR system, and the EDR retrieved from VAD scans (e,f) agrees with the vertical scan(d). The preliminary result reveal that the LiDAR observations can detect and quantify atmospheric turbulence with high spatial and temporal resolution with the VAD scan.

#### References

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